



State Key Laboratory of Numerical Modelling for Atmospheric Sciences
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Institute of Atmospheric Physics Chinese Academy of Sciences

Impacts of ENSO on the East Asian-western North Pacific monsoon

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Outline

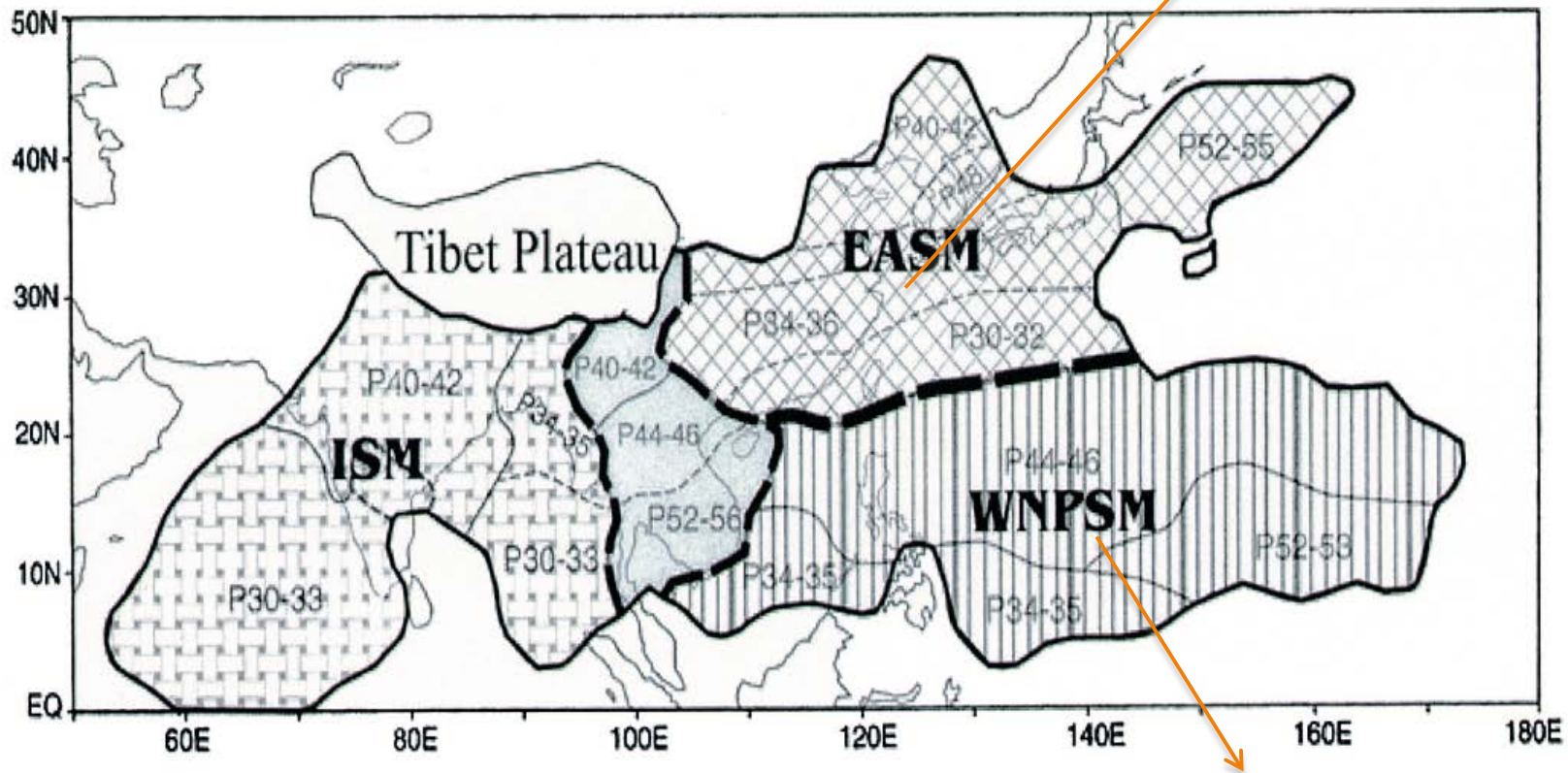


- **What is East Asian-western North Pacific (EA-WNP) monsoon?**
- **Core system linking El Nino and EA-WNP monsoon: Western North Pacific anomalous anticyclone (WNPAC)**
- **Maintenance of the WNPAC during El Nino mature winter and following spring (Remote forcing from central-eastern Pacific vs. warm pool air-sea interaction)**
- **Maintenance of the WNPAC during El Nino decaying summer (Indian Ocean basin mode vs. local cold SSTAs)**

What is EA-WNP monsoon?



Division of Asia-Pacific Monsoon



Continental monsoon, subtropical, land-ocean thermal contrast

Ocean monsoon, tropical, meridional gradient of SST

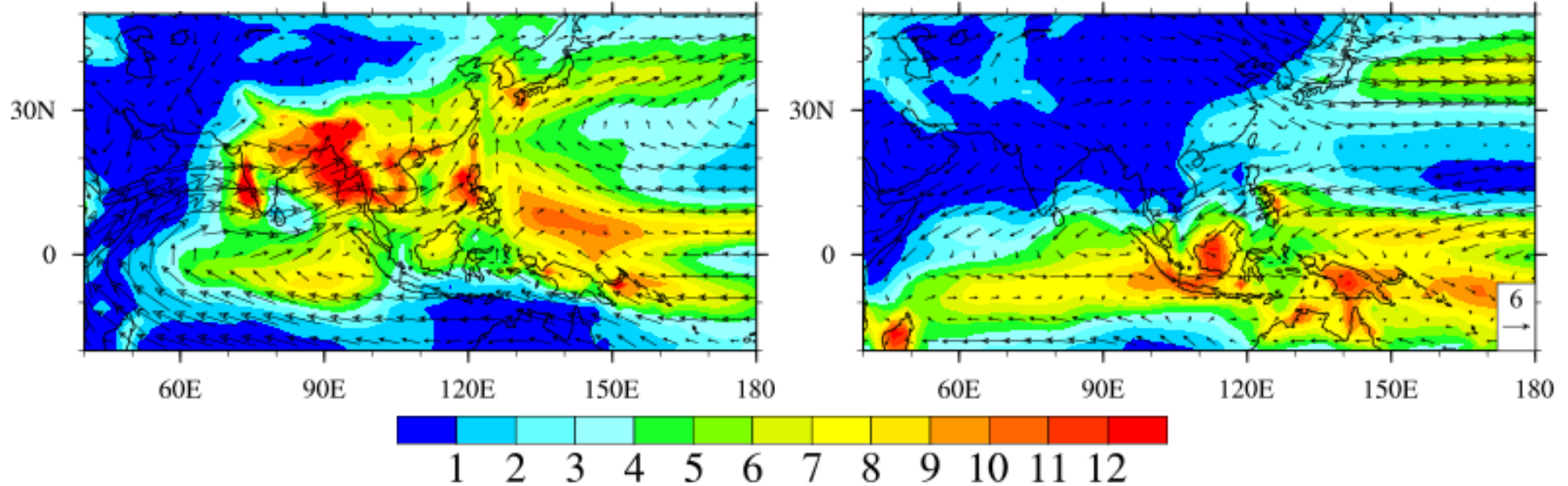


JJA

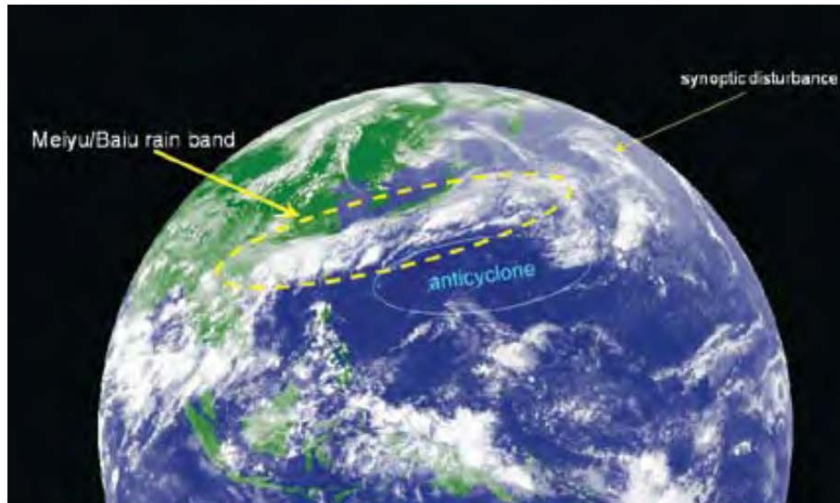
mm/day

DJF

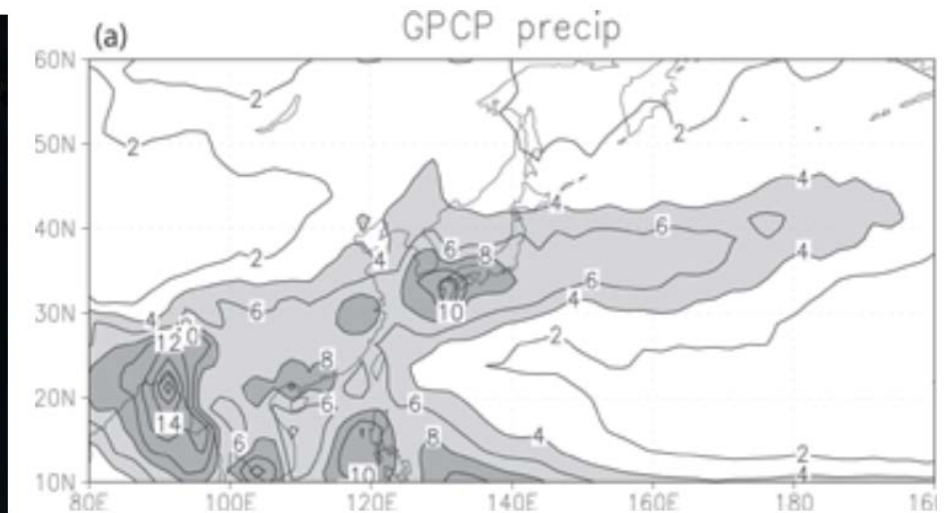
mm/day



Climatological precipitation and low-level wind



Satellite image of Meiyu-Baiu on June 12, 2008 (by Japan Meteorological Agency).



Meiyu-Baiu rainband occurs in June and early July



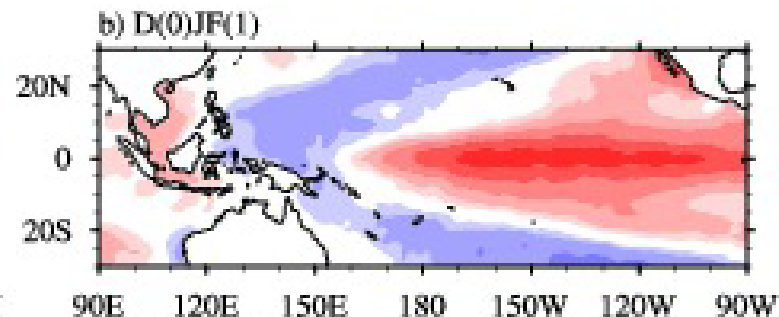
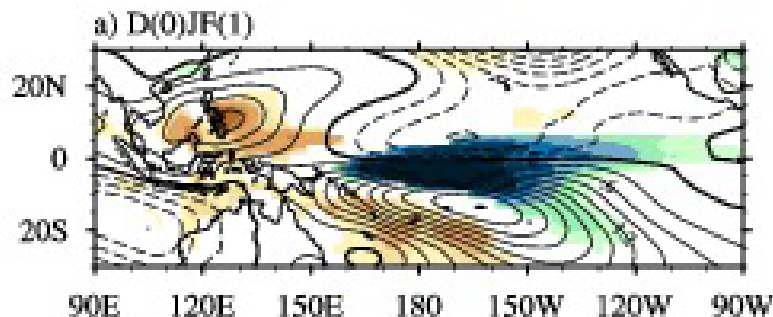
Outline



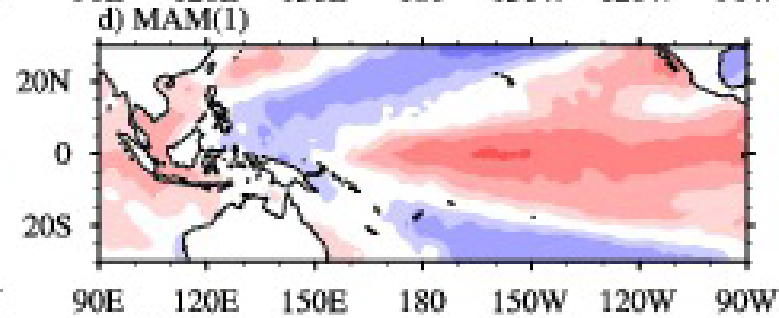
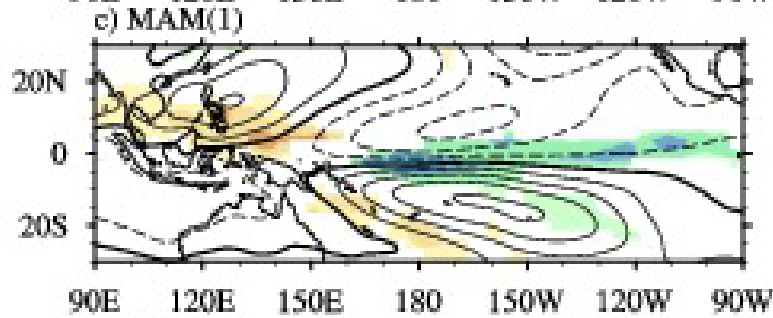
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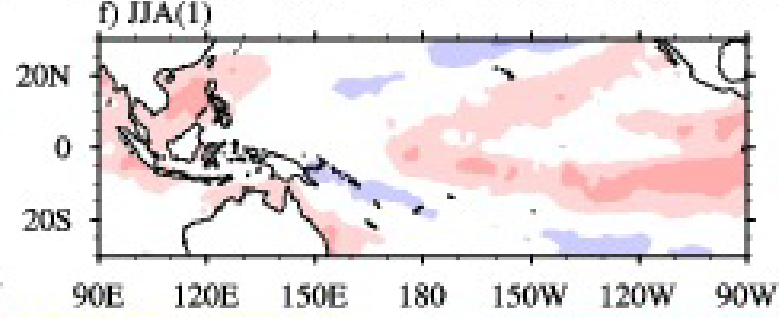
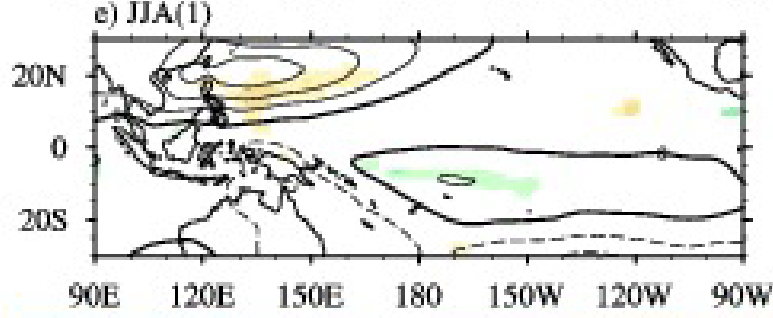
DJF

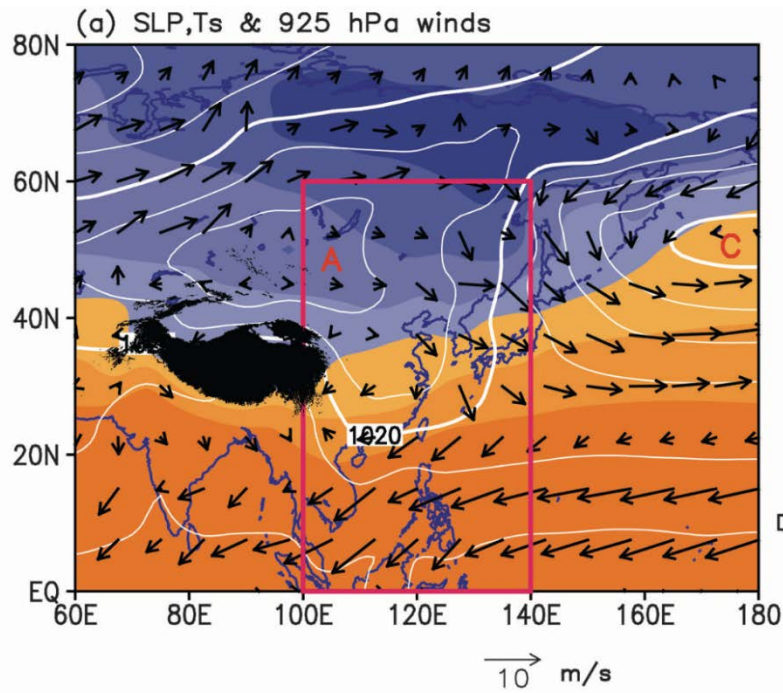


MAM

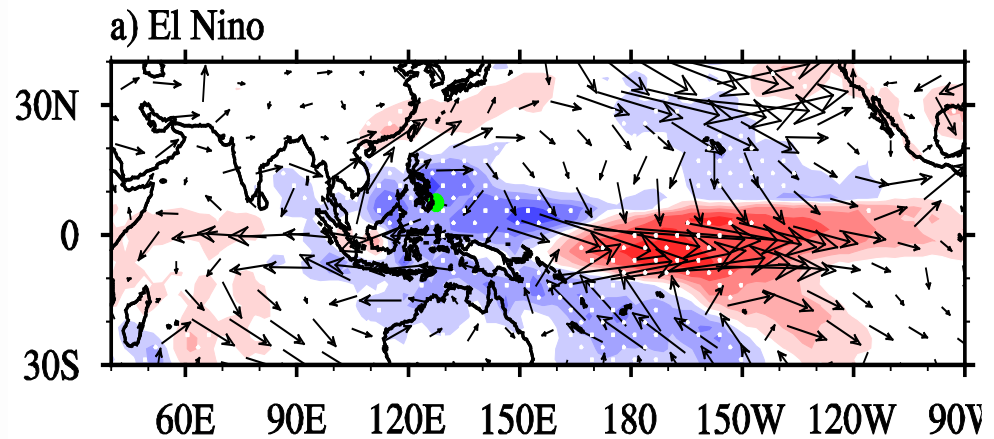


JJA

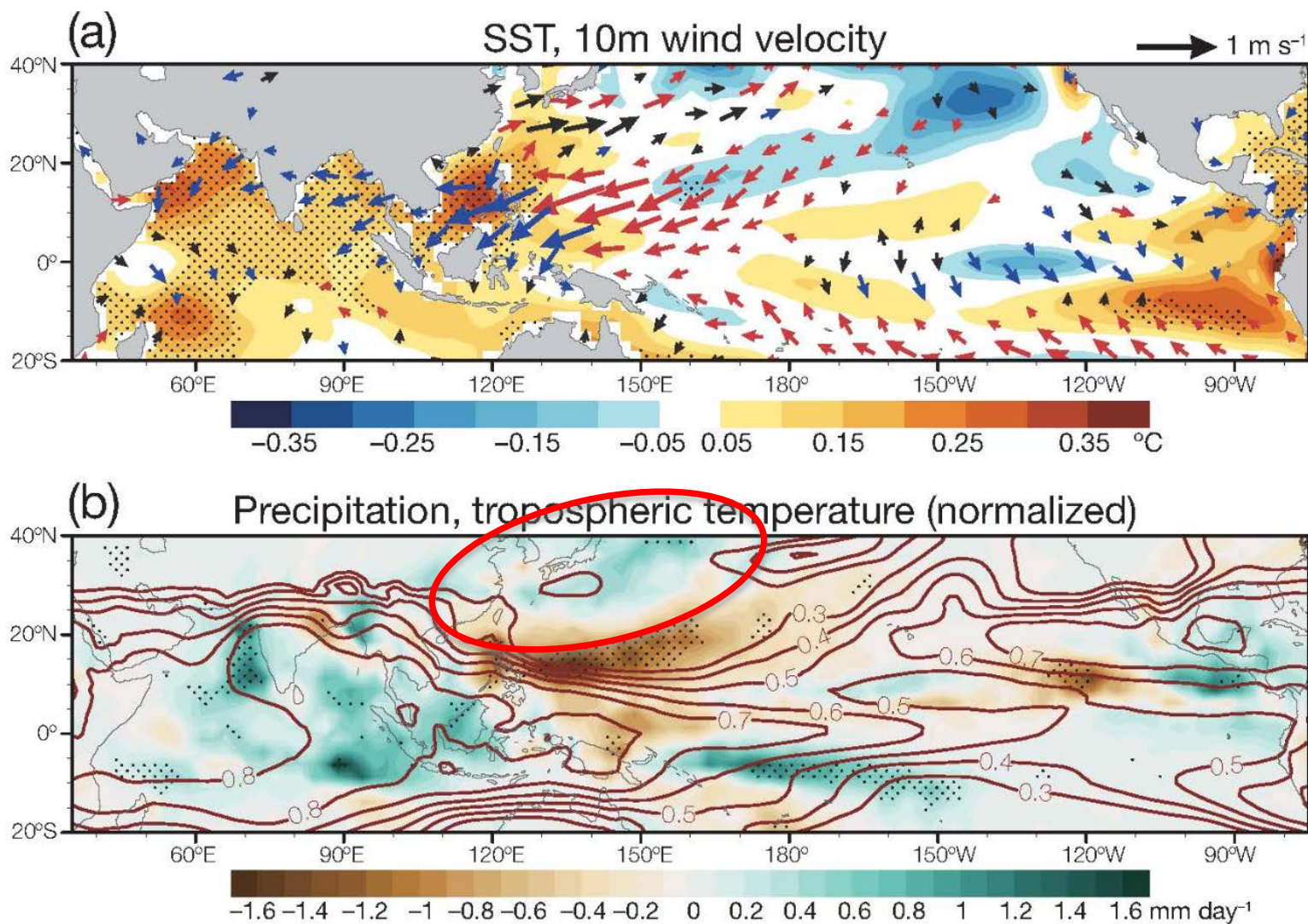




Winter climatology

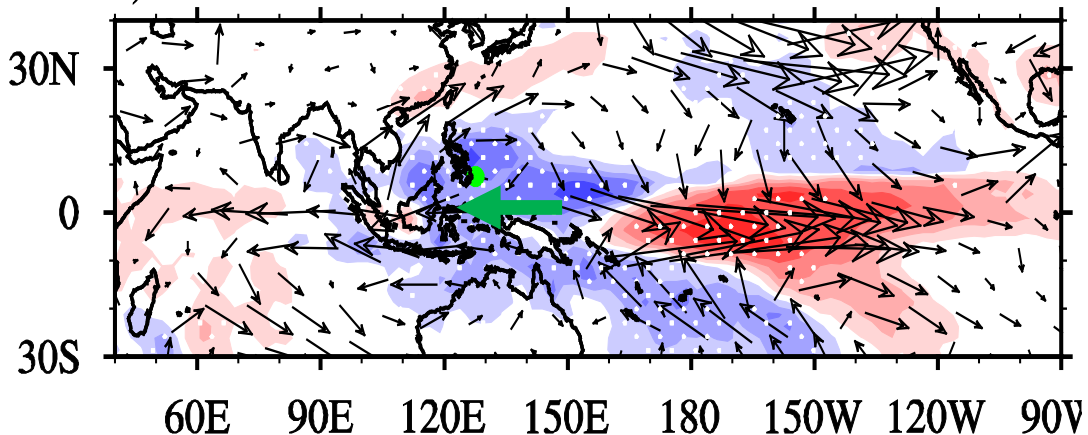


El Nino winter anomalies

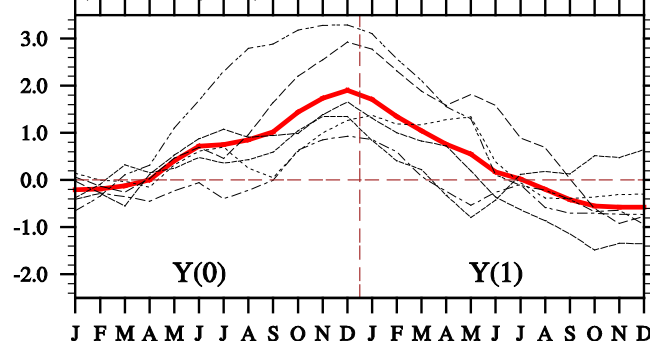




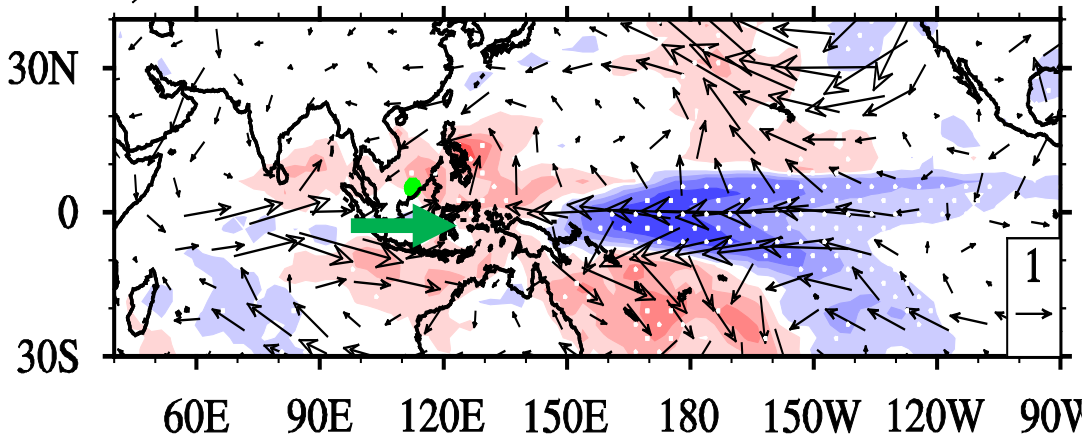
a) El Nino



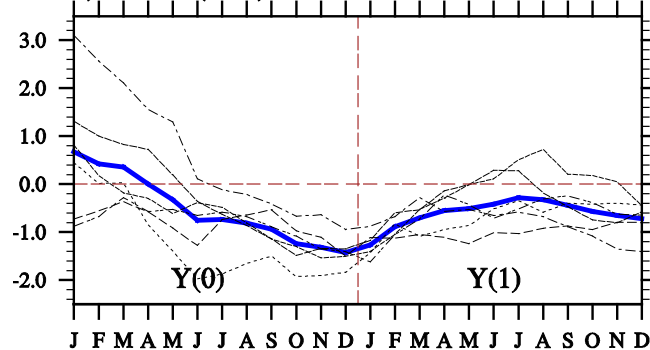
a) El Nino (OBS)



c) La Nina



b) La Nina (OBS)



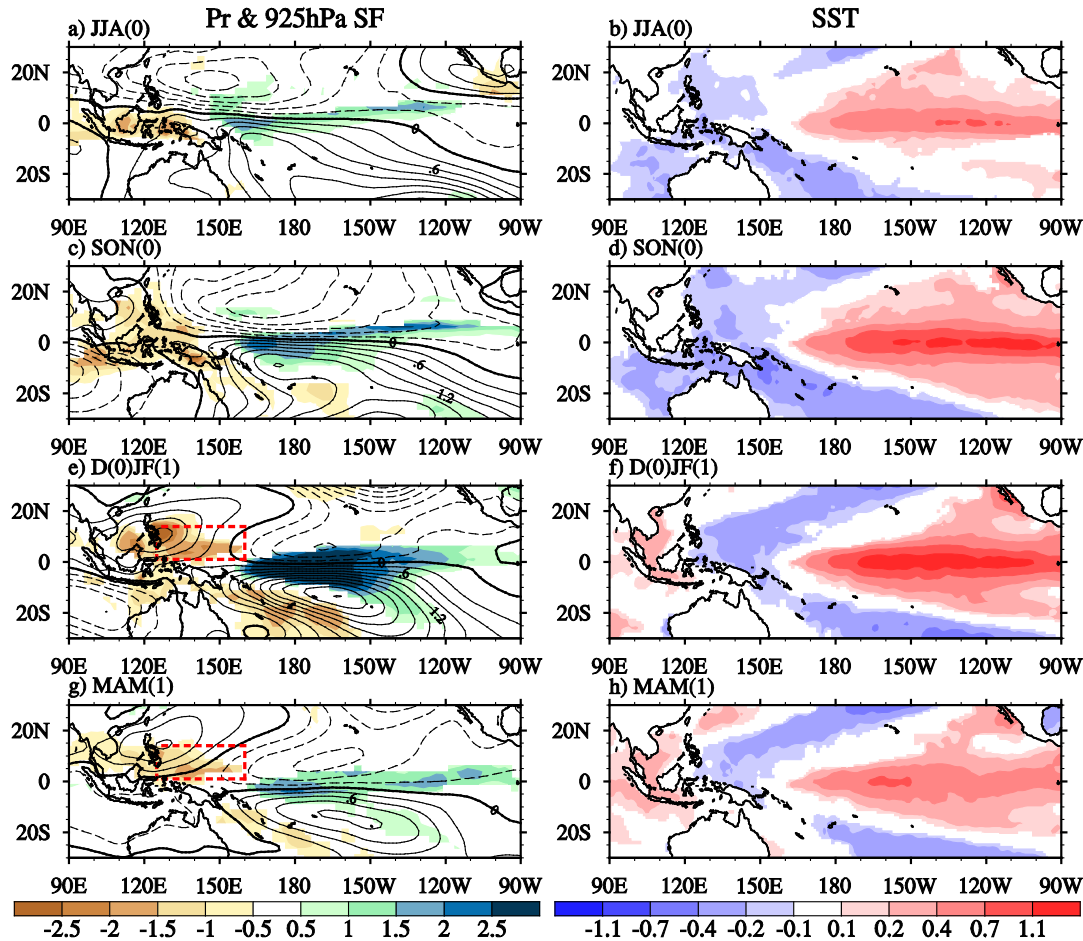
-6 -4 -2.5 -1.5 -0.5 0.5 1.5 2.5 4 6



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- **Maintenance of the WNPAC during El Nino mature winter and following spring (Remote forcing from central-eastern Pacific vs. warm pool air-sea interaction) (Wu et al. Part one and two)**
- **Maintenance of the WNPAC during El Nino decaying summer (Indian Ocean basin mode vs. local cold SSTAs)**



- What is the fundamental mechanism responsible for the maintenance of the WNPAC during El Niño mature winter and following spring?
- Why the WNPAC forms in the late fall of El Niño developing phase instead of preceding summer?



Moisture equation



$$\partial_t \langle q \rangle + \langle \mathbf{u}^{\mathbf{v}} \nabla_h q \rangle + \langle \omega \partial_p q \rangle = E - P$$

Moisture eq

$$\partial_t \langle q \rangle' + \langle \mathbf{u}^{\mathbf{v}} \nabla_h q \rangle' + \langle \omega \partial_p q \rangle' = E' - P'$$

Monthly anomalies



$$P' \approx E' - \langle \mathbf{u}^{\mathbf{v}} \nabla_h q' \rangle - \langle \mathbf{u}'^{\mathbf{v}} \nabla_h \bar{q} \rangle - \langle \bar{\omega} \partial_p q' \rangle - \langle \omega' \partial_p \bar{q} \rangle + NL$$



Moist static energy equation



$$\partial_t \langle (c_p T + L_v q) \rangle' + \langle \mathbf{u} \nabla_h (c_p T + L_v q) \rangle' + \langle \omega \partial_p h \rangle' = F_{net}'$$

MSE eq

$$F_{net}' = (S_t^\downarrow - S_t^\uparrow - R_t^\uparrow) - (-S_s^\uparrow + S_s^\downarrow - R_s^\uparrow + R_s^\downarrow - LH - SH)$$



$$\langle \omega' \partial_p \bar{h} \rangle; F_{net}' - \langle \mathbf{u} \nabla_h (c_p T + L_v q) \rangle' - \langle \mathbf{u}' \nabla_h \overline{(c_p T + L_v q)} \rangle - \langle \bar{\omega} \partial_p h' \rangle + NL$$

Monthly anomalies



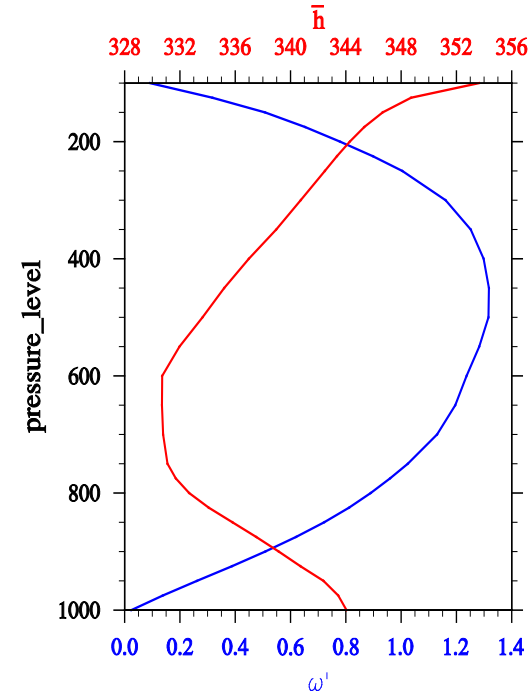
Moisture equation

$$P' ; E' - \left\langle \vec{u} \nabla_h q' \right\rangle - \left\langle \vec{u}' \nabla_h \bar{q} \right\rangle - \left\langle \bar{\omega} \partial_p q' \right\rangle - \left\langle \omega' \partial_p \bar{q} \right\rangle + NL$$

Moist static energy (MSE) equation

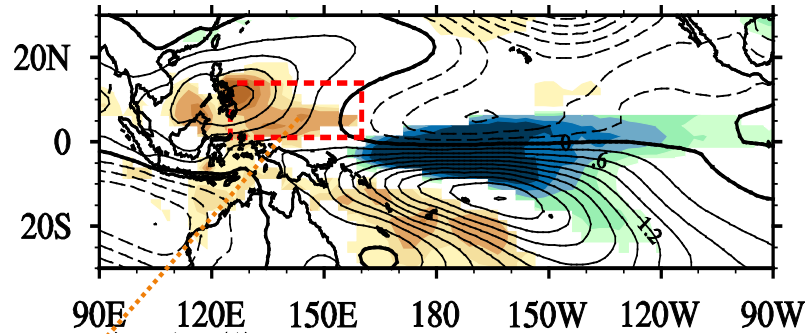
$$\left\langle \omega' \partial_p \bar{h} \right\rangle ; F_{net}' - \left\langle \vec{u} \nabla_h (c_p T + L_v q)' \right\rangle - \left\langle \vec{u}' \nabla_h \overline{(c_p T + L_v q)} \right\rangle - \left\langle \bar{\omega} \partial_p h' \right\rangle + NL$$

- In the tropics with deep convection, the vertical motion is constrained by the MSE budget balance
- If the physical processes in the right hand size of MSE equation tend to reduce the MSE in the column, descending anomalies should be generated



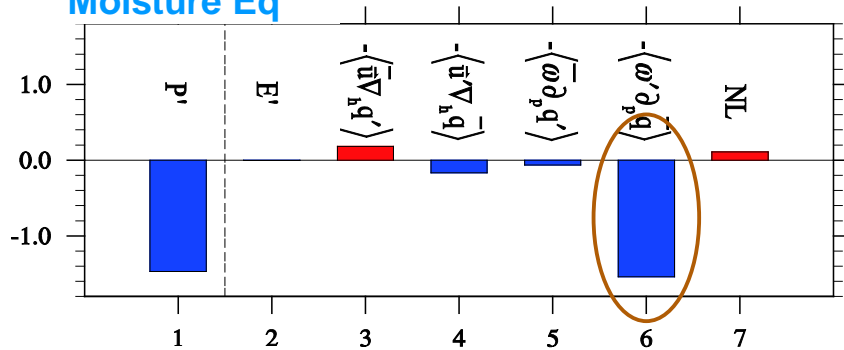


D(0)JF(1)



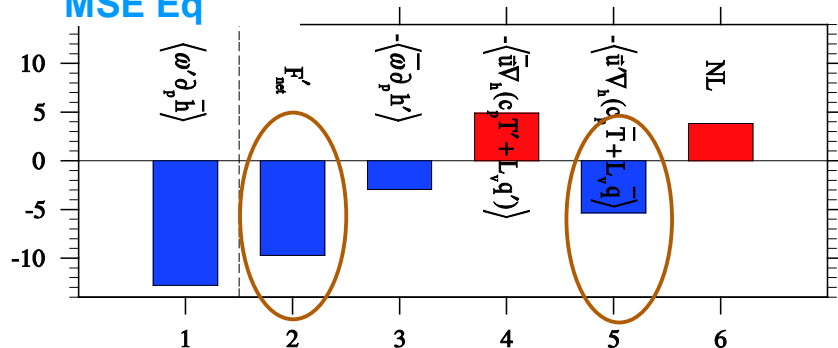
Shading: Precipitation
Contour: 925hPa stream function

Moisture Eq



$$P' ; -\langle \omega' \partial_p \bar{q} \rangle$$

MSE Eq



$$\langle \omega' \partial_p \bar{h} \rangle \sim F_{net}' - \left\langle \frac{\mathbf{v}'}{v} \nabla_h (c_p T + L_v q) \right\rangle$$

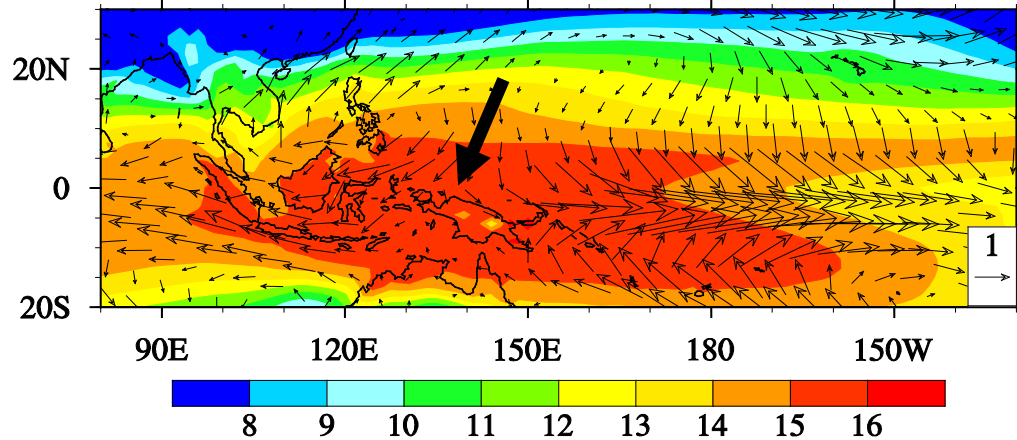
R'_{cloud}

$\langle -v' \cdot \nabla (L_v q) \rangle$

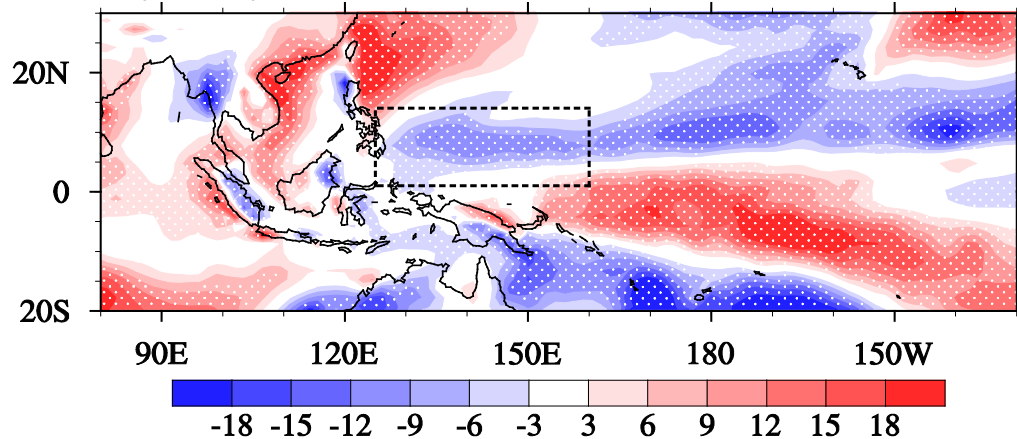


Vector: 925hPa wind anomalies

Shading: Climatological 925hPa specific humidity



b) $\langle -\bar{u}'\nabla L_v \bar{q} \rangle$

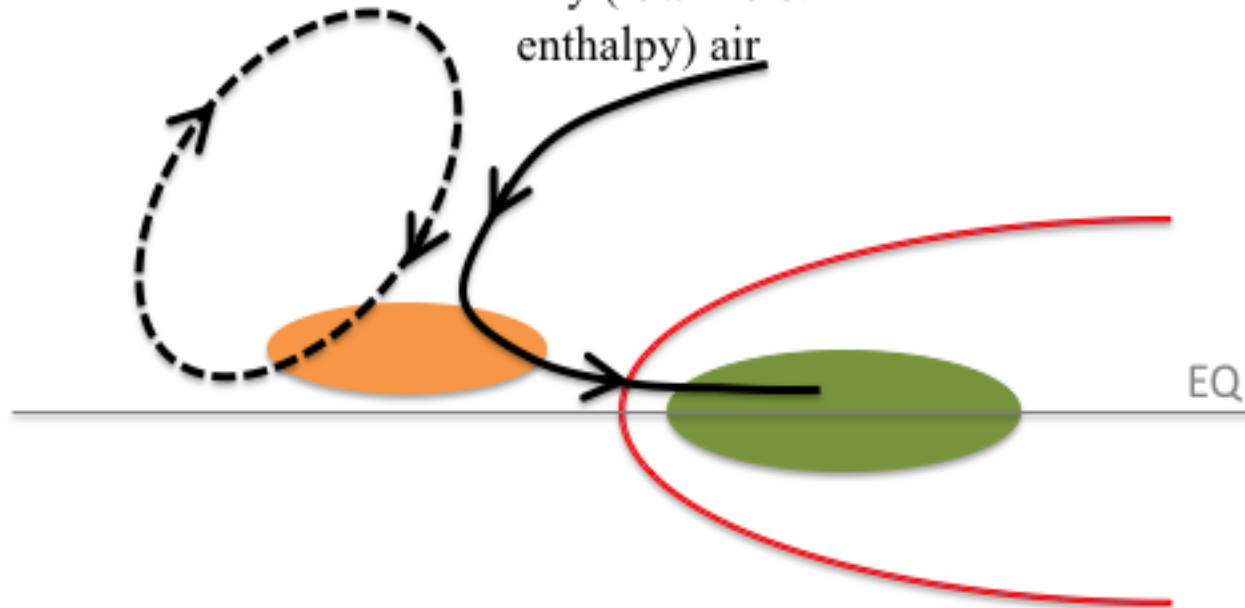


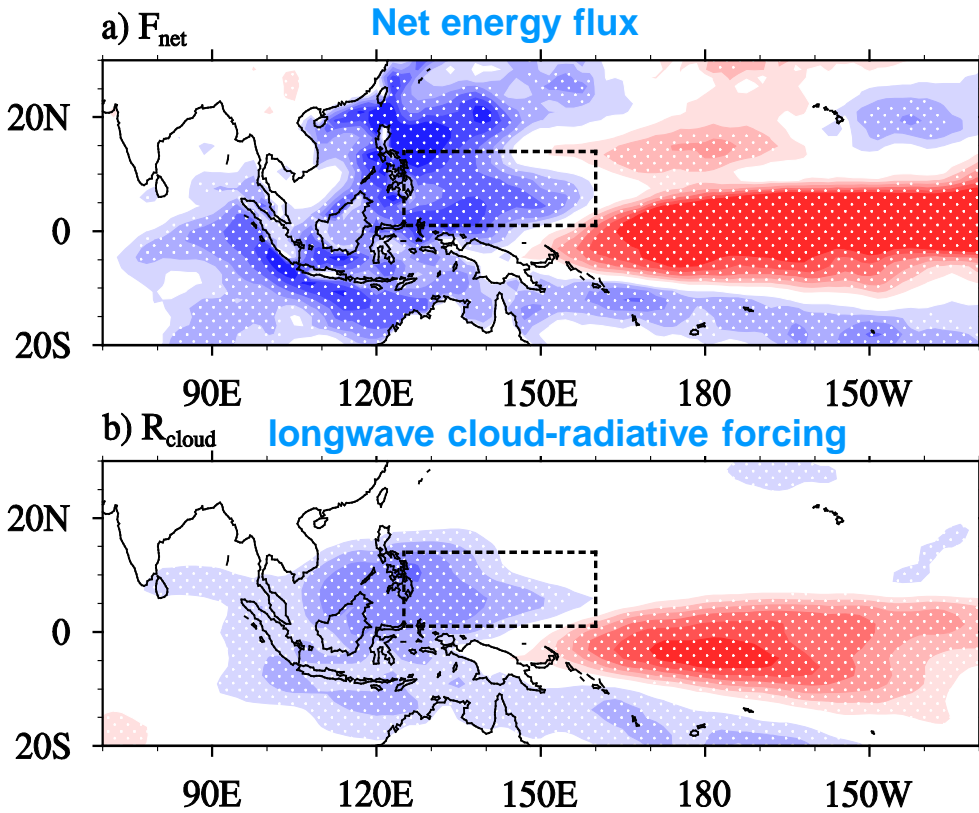
- Climatological specific humidity has negative meridional gradient
- The positive precipitation anomalies over the equatorial CEP excite twin cyclone anomalies
- The anomalous northerly component advects dry air into the tropical WNP



D(0)JF(1) and MAM(1)

Dry (low moist enthalpy) air

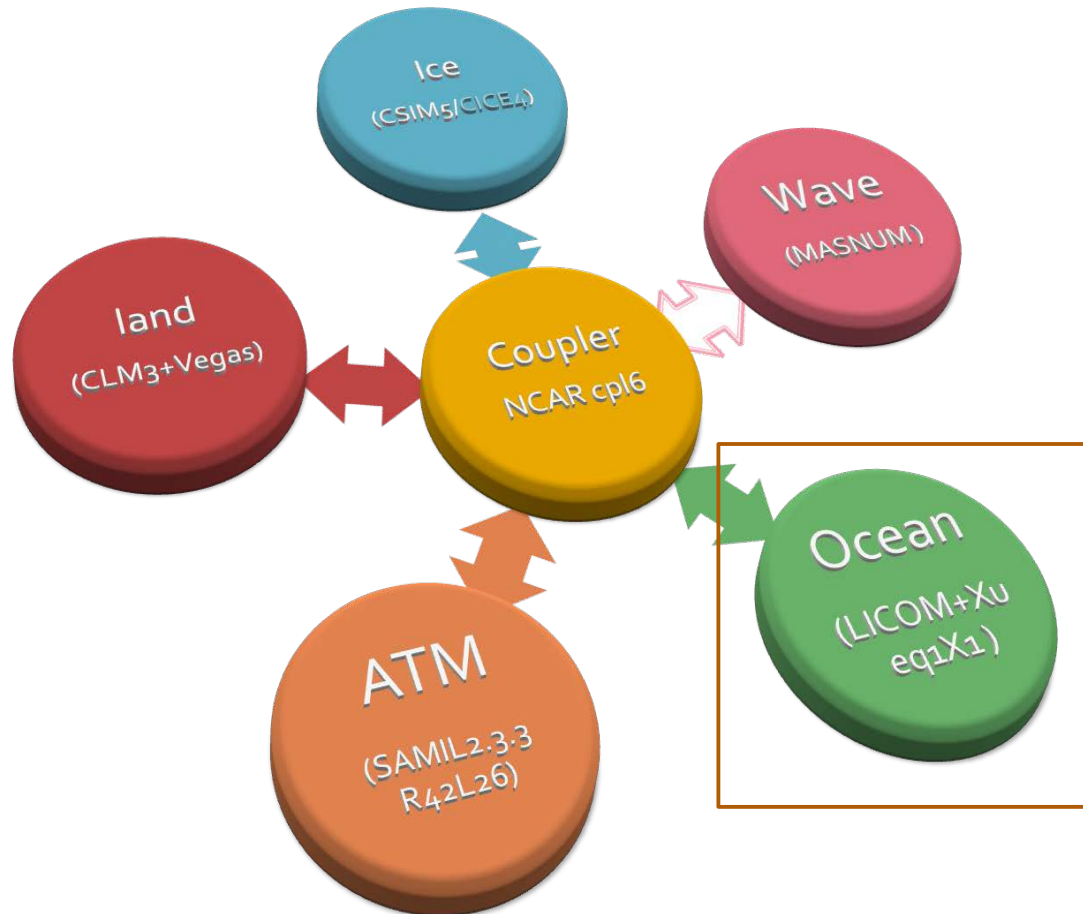




$$F_{net}' = R_{cloud}' + R_{clear}' + S_{net}' + LH' + SH'$$

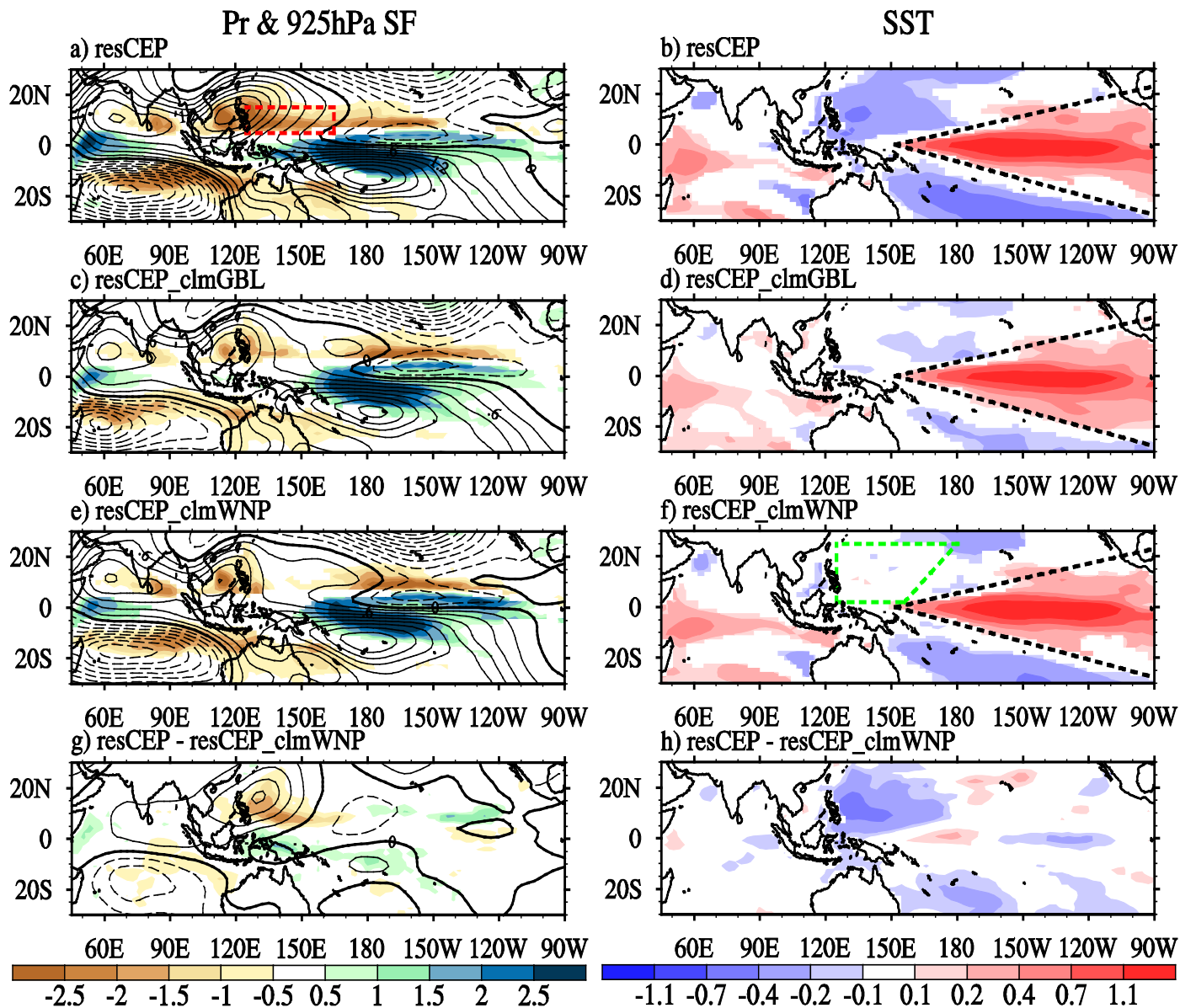
More than 60% contribution

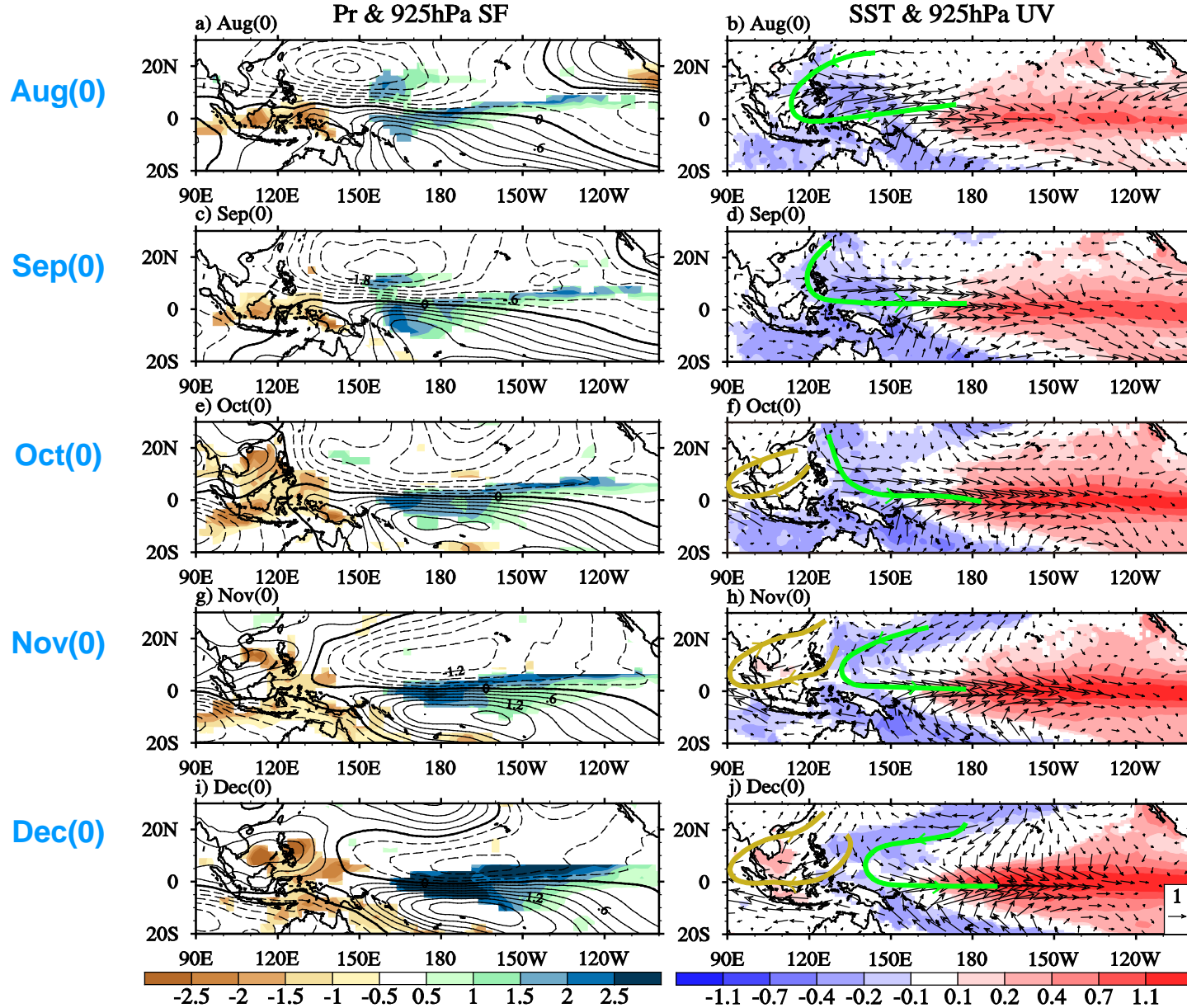
- Longwave cloud-radiative forcing anomalies are generated by an **internal positive feedback** in the tropical atmosphere
- Suppressed deep convection → decrease of deep convective cloud → Weakening of the warming effect of the longwave cloud-radiative forcing → Suppressed deep convection



Restore ocean
Temp in equatorial
CEP to OBS

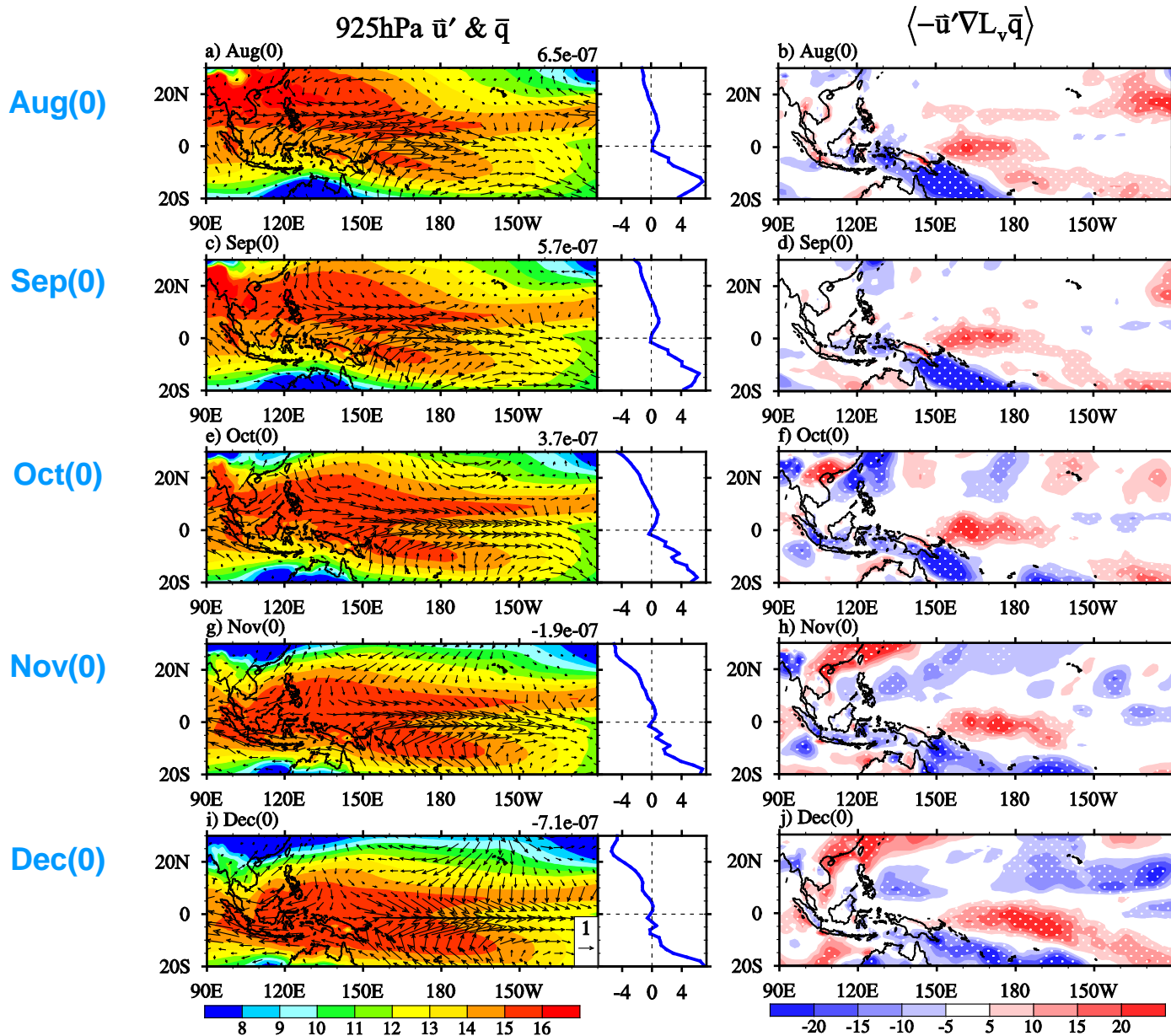
FGOALS-s2 is a Coupled GCM





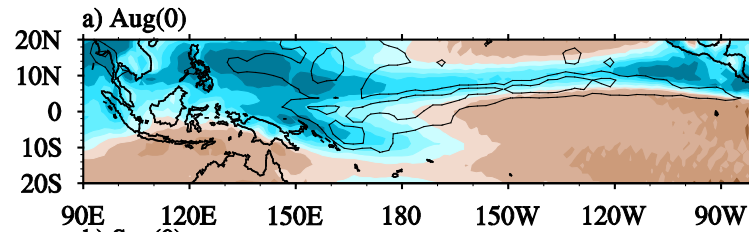


Temporal evolution of anomalous advection of moist enthalpy

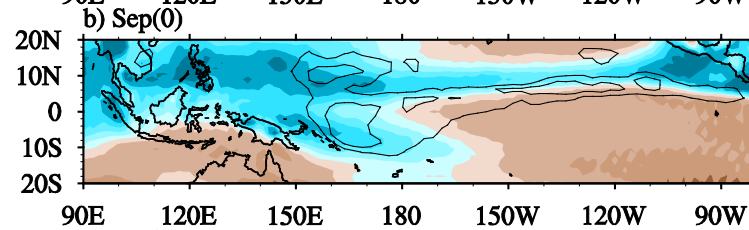




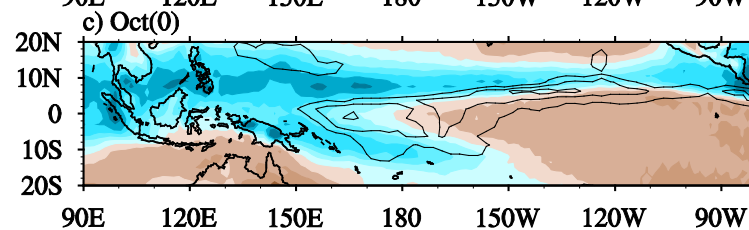
Aug(0)



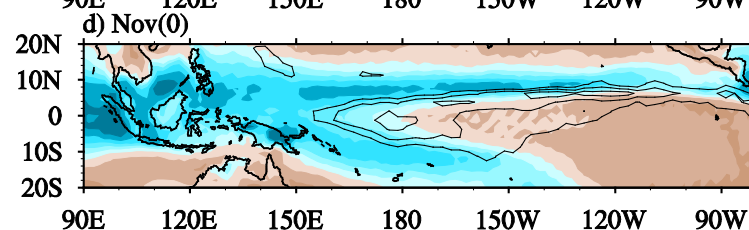
Sep(0)



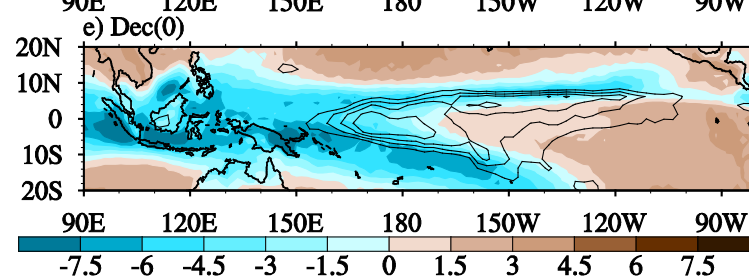
Oct(0)

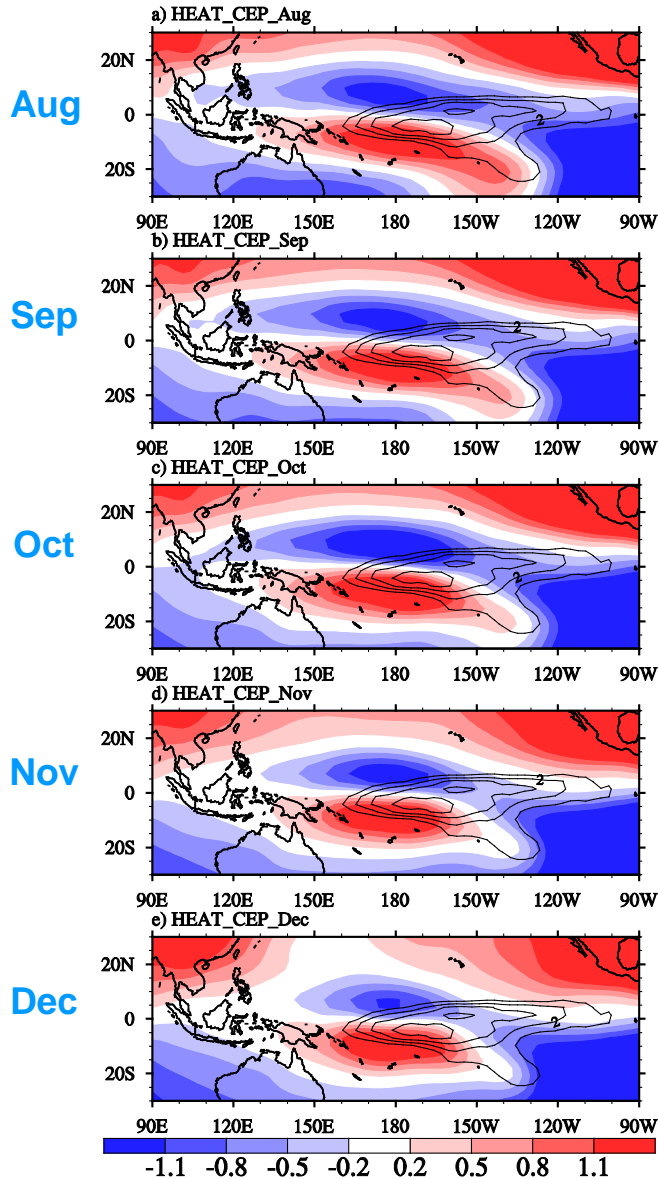


Nov(0)



Dec(0)

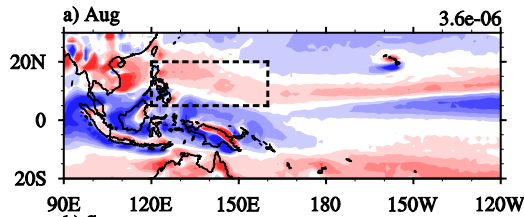




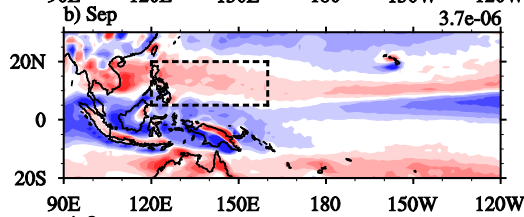
- Constructed the heating field using the pattern of the positive precipitation anomalies over the equatorial CEP during the El Niño mature winter.
- The heating field was used to drive the model, with the background mean states specified as the climatology of August, September, October, November and December derived from the reanalysis data



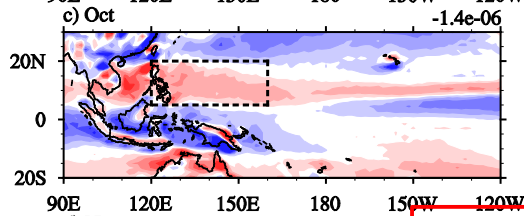
Aug



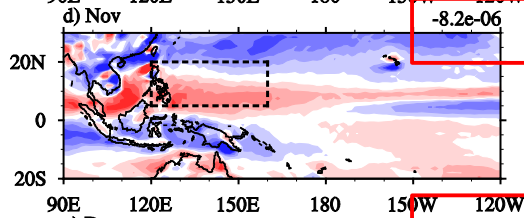
Sep



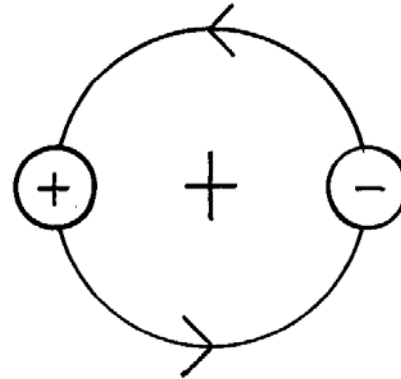
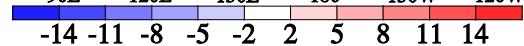
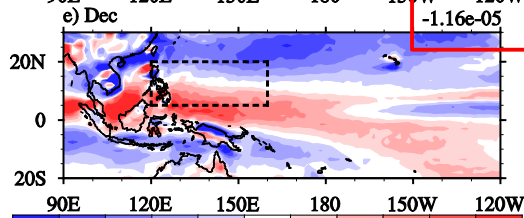
Oct



Nov



Dec



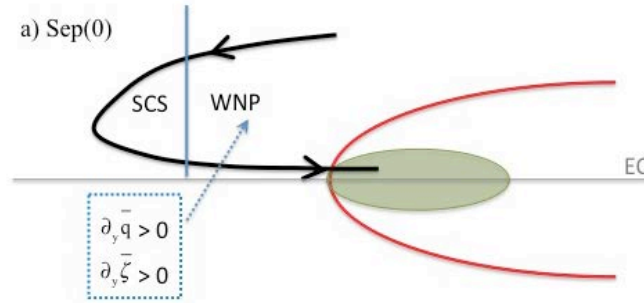
From Hoskins and Ambrizzi 1993

- The meridional gradient of the vorticity over the WNP transforms from positive to negative.
- The sign change in late fall increasingly offsets the beta effect and thus reduces the **westward stretch of the Rossby-wave gyre anomalies**.
- The northern branch of the twin cyclonic anomalies induced by the El Niño heating withdraws eastward, leaving space for the formation of the WNPAC.

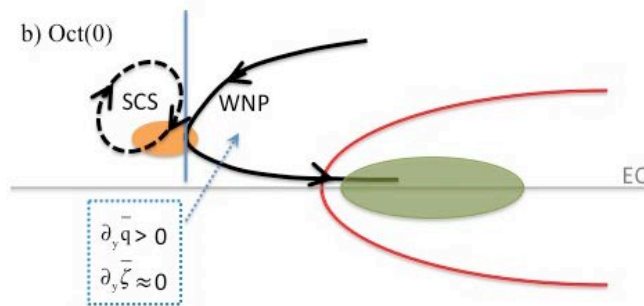
climatological relative vorticity



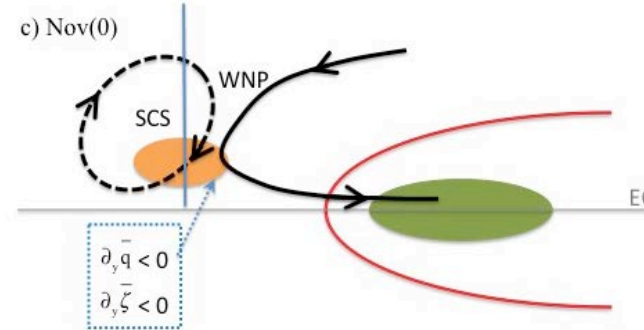
Sep(0)



Oct(0)



Nov(0)

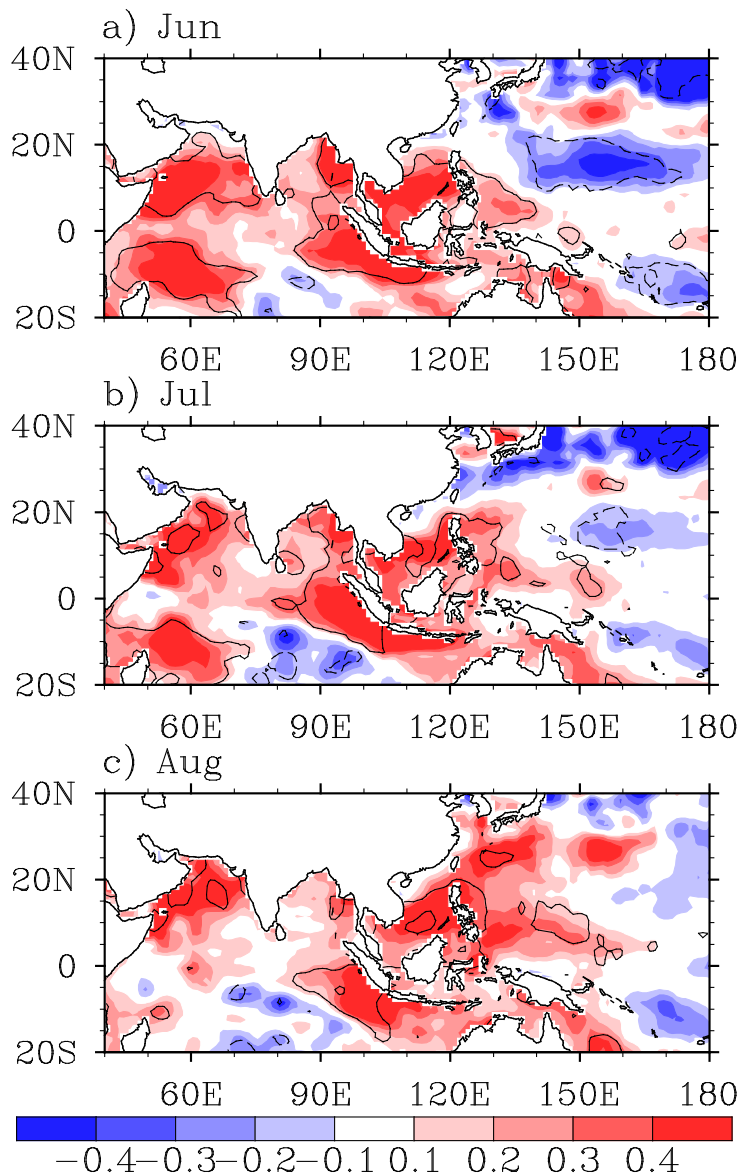




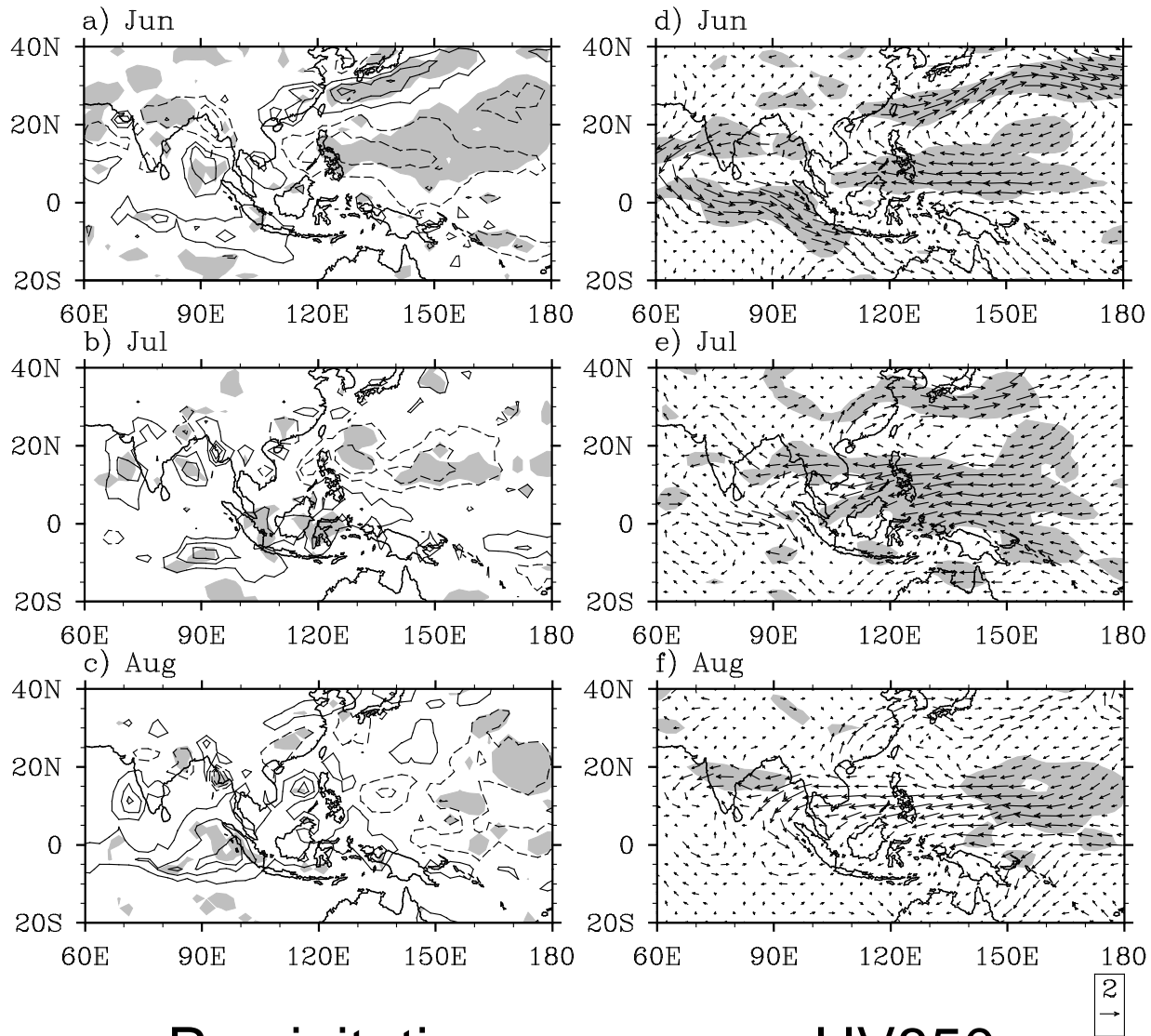
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- Cold SSTA in the WNP gradually decay and withdraw eastward from June to August.
- Indian Ocean basin mode (IOBM) maintains throughout the summer.

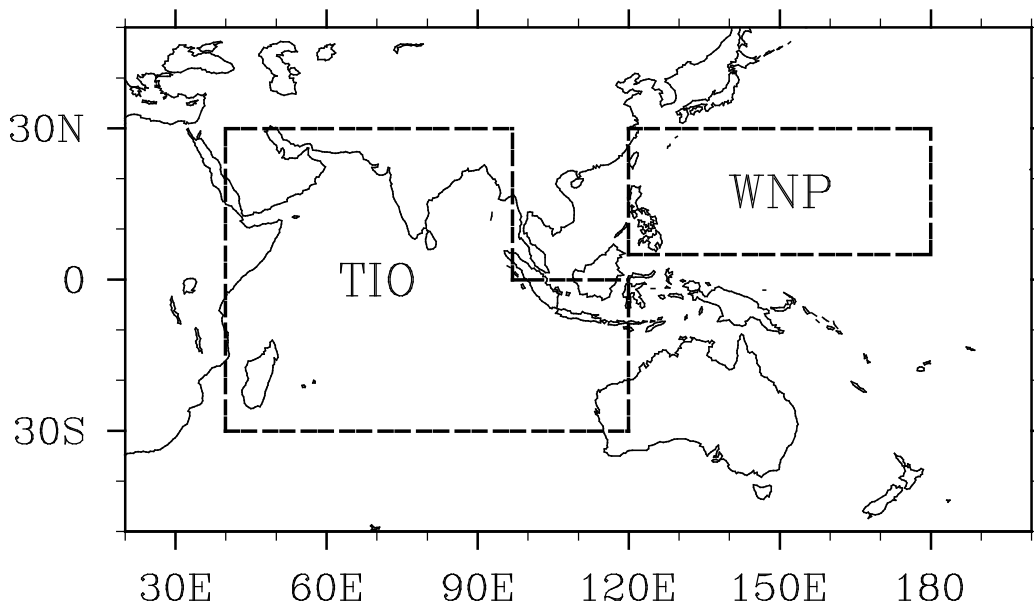


Precipitation

UV850



Idealized numerical experiments by AGCM



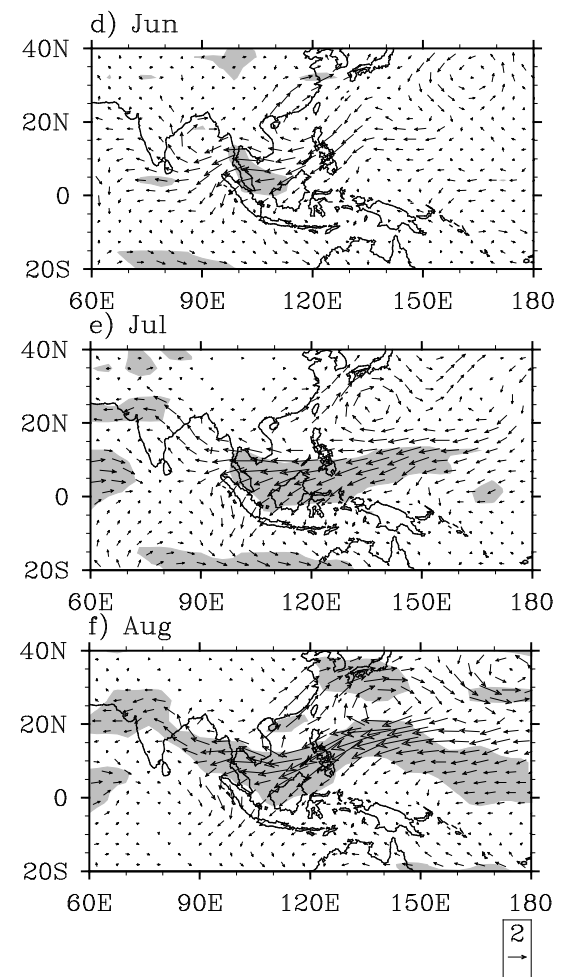
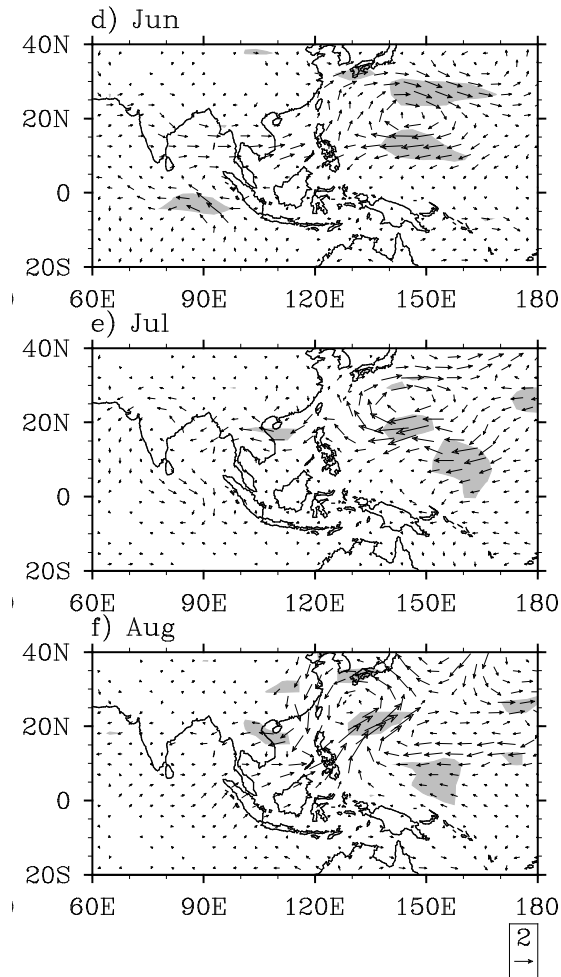
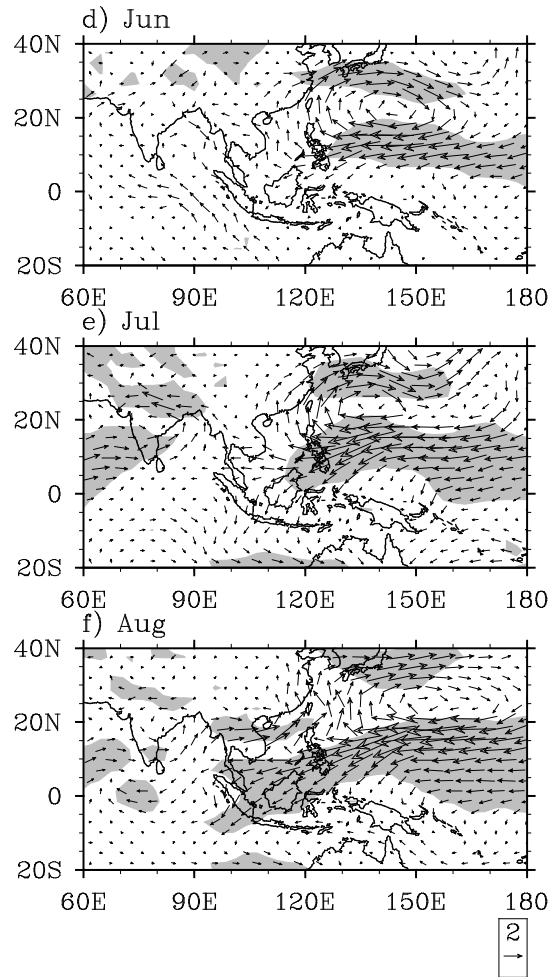
Experiment	SST forcing	Integration
Control run (CTRL)	Global climatological SST	20 years
Global forcing (GB)	Add SSTA to climatological SST in global ocean	20 realizations
Indian Ocean forcing (TIO)	Add SSTA to climatological SST in the tropical Indian Ocean only	20 realizations
Western North Pacific forcing (WNP)	Add SSTA to climatological SST in the western North Pacific only	20 realizations

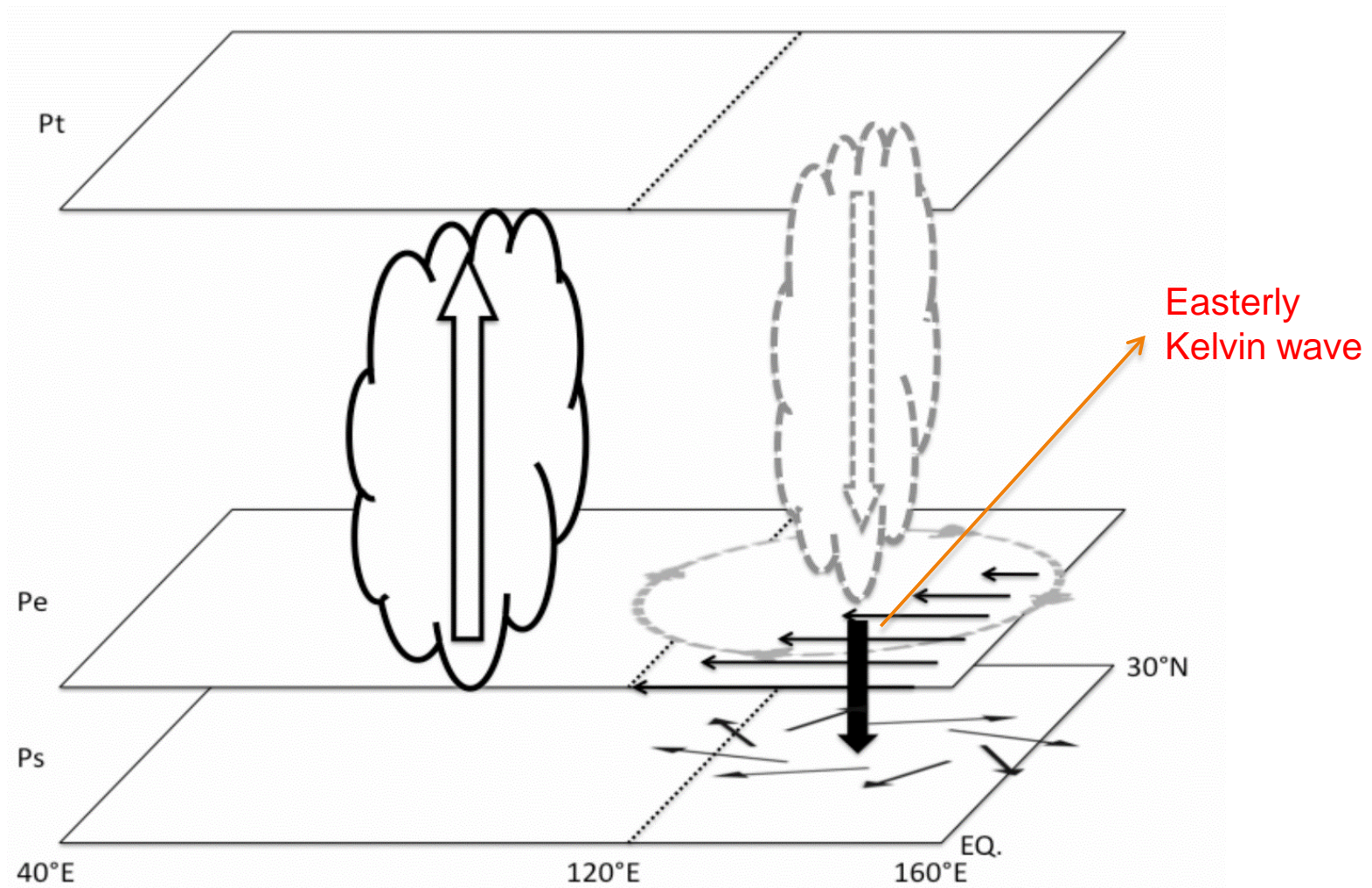


GB

WNP

TIO







Summary



- WNPAC plays a central role in linking El Nino and EA-WNP monsoon
- WNPAC maintains from El Nino developing winter to decaying summer
- WNPAC is driven by El Nino remote forcing through moist enthalpy advection mechanism during El Nino developing winter and following spring
- WNPAC is driven by Indian Ocean basin wide warming and local cold SSTAs during El Nino decaying summer



Reference



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- Xie, S.-P., et al., 2016: Indo-western Pacific ocean capacitor and coherent climate anomalies in post-ENSO summer: A review. *Adv. Atmos. Sci.*, 33, 411-432
- Wu B., et al., 2017: Atmospheric dynamic and thermodynamic processes driving the western North Pacific anomalous anticyclone during El Niño. Part I: Maintenance mechanisms. *Journal of Climate*, In press. doi.org/10.1175/JCLI-D-16-0489.1
- Wu B., et al., 2017: Atmospheric dynamic and thermodynamic processes driving the western North Pacific anomalous anticyclone during El Niño. Part II: Formation processes. *Journal of Climate*, In press. doi.org/10.1175/JCLI-D-16-0495.1
- Wu, B. et al., 2010a: Asymmetry of Atmospheric Circulation Anomalies over the Western North Pacific between El Niño and La Niña. *J. Climate*, 23, 4807-4822.
- Wu, B., et al., 2010b: Relative contributions of the Indian Ocean and local SST anomalies to the maintenance of the western North Pacific anomalous anticyclone during El Niño decaying summer. *J. Climate*, 23, 2974-2986
- Wu, B., et al. 2009, Seasonally Evolving Dominant Interannual Variability Modes of East Asian Climate, *Journal of Climate*, 22, 2992-3005
- Li and Wang, 2005: A Review on the Western North Pacific Monsoon: Synoptic-to-Interannual Variabilities
- Wang, B., and Lin H., 2002: Rainy seasons of the Asian-Pacific monsoon. *J. Climate*, 15, 386- 398.

A blue-tinted globe of the Earth is centered in the background. The text "Thank You!" is overlaid in the center of the globe. The text is white with a pink-to-white gradient and a black outline. The background transitions from blue at the top to orange at the bottom.

Thank You!